

STABILITY OF NATURAL SLOPES
IN THE MACKENZIE VALLEY



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STABILITY OF NATURAL SLOPES IN THE MACKENZIE VALLEY



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THE STABILITY OF NATURAL SLOPES IN THE MACKENZIE VALLEY

by

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Geological Survey of Canada

Department of Energy, Mines
and Resources

for the

Environmental-Social Program
Northern Pipelines

July 1973

Canada

Environmental-Social Committee
Northern Pipelines,
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Report No. 73-9

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10 maps



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The data for this report were obtained as a result of investigations carried out under the Environmental-Social Program, Northern Pipelines, of the Task Force on Northern Oil Development, Government of Canada. While the studies and investigations were initiated to provide information necessary for the assessment of pipeline proposals, the knowledge gained is equally useful in planning and assessing highways and other development projects.

TABLE OF CONTENTS

	Page
Summary	1
Introduction	2
Review of Current Knowledge	2
Classification System	3
Slope Erosion - General	3
Slope Stability in the Mackenzie Valley	4
Interpretation of Field Investigations	5
Engineering Significance of Mass Wasting in the Mackenzie Valley	6
Requirements for Further Study	7
Bibliography	8
Figures	9
Table 1 River bank stability maps. Legend	18

FIGURES

Illustrating types of slope failures found in the
Mackenzie Valley.

Page

Figure 1	Earthflow, elongate (Qa) Mackenzie River, north of Wrighley.	9
Figure 2	Earthflow, lobate (Qa) Mackenzie River, north of Fort Simpson.	9
Figure 3	Earthflows in recently burnt area west of Wrighley (Qa).	10
Figure 4	Closer view of Figure 3.	10
Figure 5	Detachment slides in granular soil (Qa) Mackenzie River near Blackwater River.	11
Figure 6	Detachment slide in Cretaceous shale (Ka).	11
Figure 7	Large scale retrogressive slide; (Q1) Mackenzie River near Root River.	12
Figure 8	Large scale retrogressive slide; (Q1) Mackenzie River near Root River.	12
Figure 9	Large scale retrogressive failure; Quaternary sediments (Q1) Mackenzie River opposite Fort Norman.	13
Figure 10	Large scale retrogressive slides, Quaternary sediments (Q1) Mackenzie River south of Fort Norman.	13
Figure 11	Large scale retrogressive failure in Cretaceous shale (K1) Arctic Red River.	14
Figure 12	Large scale retrogressive failure in Cretaceous shale (K1) Mountain River.	14
Figure 13	Large scale slumping in Tertiary sediments (Te) Mackenzie River south of Fort Norman.	15
Figure 14	Undercutting of detached frozen slumped block (Q1) Mackenzie River north of Sans Sault Rapids.	15
Figure 15	River bank in Devonian sediments (D) Liard River.	16
Figure 16	Slump in Devonian sediments (D) Liard River.	16
Figure 17	Large scale multiple retrogressive slide (Q1) Mackenzie River north of Fort Simpson.	17
Figure 18	Gullying in glacial till (Qa) Mackenzie River south of Fort Simpson.	17

SUMMARY

A program of mapping stability of river banks of the Mackenzie River and its tributaries within the adjacent plains region was carried out in ten map areas between Fort Simpson and Fort Good Hope, N.W.T. Mapping was done according to an engineering-geological classification system which relates certain conditions of stability and topography to the geology. In general, slope failures are of two main types - shallow active layer failures, and larger scale multiple retrogressive slides characteristic of high banks of Quaternary sediments and Cretaceous Shales.

As potential river crossing sites are evaluated for pipelines or roads, factors to be considered should include the following:

- the retrogressive nature of the larger types of slides
- the abrading action of river ice on river banks and bottom during break-up
- the adverse effects of forest fire on stability of slopes
- the effects of ice jamming in raising upstream water and ice levels

INTRODUCTION

A rational approach to the planning, design and construction of transportation systems in the Mackenzie Valley must take into account the mass wasting of sloping terrain, particularly of slopes forming the sides of river valleys. The overall objective of this project is to provide certain geologically based information which will assist in efforts to minimize the environmental disturbance associated with future construction on, and adjacent to, such slopes.

The information presented on the appended map sheets and in the text represents the results of a field mapping operation initiated during the summer of 1971 and continued in 1972. The maps provide an inventory of the river banks of the Mackenzie River and its main tributaries within the adjacent plains region. They are classified according to a system, shown in the legend, which relates types of failures to geologic and topographic factors. Helicopter borne field mapping was supplemented with ground verification as required. Field mapping was done on high level air photos and 1:50,000 scale hydrographic charts. The information was reduced and compiled on the appended 1:250,000 scale topographic sheets. The following map areas were covered:

95 H, I, J, O, N

96 C, D, E

106 H, I

In conjunction with the inventory-classification aspect of the mapping program, a qualitative assessment has been made of the geomorphic processes and the failure mechanics associated with mass wasting. In addition, the significance of river bank erosion as it applies to engineering activities has been considered.

REVIEW OF CURRENT KNOWLEDGE

Because of the hazards often associated with landslides and slope failures, considerable effort has been expended in exploring the causes of these phenomena and in the prevention and remedial treatment of them. Analytical methods for determining the stability of slopes in terms of factors of safety form an integral part of the discipline of soil mechanics. Most of these techniques have been developed for application to unfrozen soils, less being known of the failure mechanics associated with mass movements in permafrost terrain. The recovery and testing of undisturbed borehole samples of frozen soil and monitoring of pore pressure conditions are required in order to do such analyses and costs of these operations are considerably increased because of the isolated conditions.

Recent studies of failure mechanics at selected sites along the Mackenzie River include those initiated in 1972 by Dr. N. Morgenstern and Mr. E. MacRoberts of the University of Alberta.

CLASSIFICATION SYSTEM

A number of classification systems have been devised in association with studies of landslides and related phenomena. In general such systems have their roots either in the earth sciences or the related engineering discipline of soil mechanics. The classification system used in this study attempts to straddle the boundary between geology and engineering and is tailored to the reconnaissance - inventory nature of the mapping operation to which it was applied.

The system is intended to serve the practical purposes of engineering and construction, and for this reason the different scales of mass movement are given emphasis. This approach is facilitated by the two orders of magnitude of material displaced as slopes fail. The smaller failures are confined to the shallow active layer, the top few feet of annually thawed ground overlying permafrost. The much larger slides or slumps, and in some cases flows, involve displacement of an entire river bank which may be 100 feet or more in height. In such cases, frozen as well as thawed soil and rock may be displaced. The magnitude of the environmental hazard bears a parallel relationship to the scale or size of failure. The primary subdivision is made in terms of the geology since a relationship exists between mode of failure, strength properties of earth materials, and the age of deposition. Hence a degree of predictability of the behaviour of slopes, can be established if the geologic age and composition of the soil or rock in each area are known.

This classification system (see Table 1, page 18) is a revision of one devised to describe slope failures in the Fort Good Hope - Sans Sault Rapids area (Isaacs and Code, 1972).

SLOPE EROSION - GENERAL

Retreat of river valley slopes occurs as a result of geomorphic processes that can be categorized either as mass movement or mass transport. Mass movement refers to the downslope movement of earth materials under the influence of gravity without the assistance of a transporting media. Mass transport, on the other hand, refers to the transport of fragments of earth materials suspended in a transporting medium, either water, ice or wind. These two types of processes do not necessarily act independently, and may be in effect either at the same time or sequentially.

In the area mapped, mass movement is expressed in the following terms:

- falls: rock falls and soil falls
- slides: slumps, translational slides, detachment slides
- flows: earthflows, solifluction

The meanings intended by this terminology are explained below:

(i) The terms "rock fall" and "soil fall" are almost self-explanatory and indicate free fall motion of fragments with some related rolling or sliding.

(ii) "Slides" are types of mass movement in which the detached mass is not significantly deformed while in motion. They are initiated by failure either in shear or tension along a finite failure plane. "Slumping" refers to rotational failure in shear with a backtilting of the displaced unit. "Translational sliding" refers to movement along a planar surface which in permafrost terrain is probably an unfrozen clay or silt layer within or below the permafrost layer. The term "detachment slide" (O.L. Hughes, 1972) refers to downslope displacement of the thawed active layer of organic material, thawed soil, or weathered shale. "Earthflow" refers to the downslope displacement of earth materials which behave as viscous fluids during motion. Very rapid earthflows are sometimes referred to as mudflows. The term solifluction has been defined as "slow flowing from higher to lower ground of masses of waste saturated with water". It refers to a periglacial phenomenon, and rates of movement are generally in the order of 2-5 cm per year (Hamelin and Cook, 1967). "Slopewash" refers to the downslope mass transport of materials suspended in water without the organization of drainage into a system of rills and gullies. "Gully erosion" or "gullying" is a similar phenomenon but one in which the water and suspended materials are carried in channels.

SLOPE STABILITY IN THE MACKENZIE VALLEY

In any stability analysis of slopes, either natural or man-made, certain factors or conditions must be taken into account. These are (a) the strength properties of the materials (b) hydrological conditions (c) topography (d) active erosional processes which may alter the stability.

In periglacial regions factors which relate to the presence of ground ice and to the severe climatic conditions should also be considered.

(i) The presence of the impermeable permafrost layer just below the ground surface causes surface water to drain laterally rather than being partially absorbed into the ground as in temperate climates. Hence higher moisture contents are found in the active layer than are found in equivalent surface layers of non-permanent terrain. These higher water contents result in a lowering of shear strengths of the soils, thus reducing their overall stability on inclined terrain.

- (ii) River ice carried downstream during spring break-up, abrades the river banks at various levels, removing vegetative cover, steepening and unloading the toe of slopes thus generally reducing the stability.
- (iii) Due to the presence of ground ice, moisture contents in fine grained (cohesive) soils are often higher than the liquid limit. Such soils when thawed become highly mobile.
- (iv) Slumped or sliding blocks of frozen ground usually remain relatively undeformed during the period of movement. Deformation would proceed if thawing were to occur after failure.
- (v) Within certain limits, the shear strength of soil in the permafrost layer increases in some proportion to the amount of ground ice present, and in proportion to decreasing ground temperature.
- (vi) Melting ice wedges adjacent to slopes may provide planes of weakness which then develop into tension cracks or channels capable of capturing surface run-off and initiating gullyng.
- (vii) Regeneration of burnt organic cover is slower than in more temperate climates. This slows the rate of stabilization of slopes failing because of the loss of organic cover.
- (viii) Ice jamming in rivers during spring break-up has a number of effects which tend to reduce the stability of adjacent slopes. These effects are:
 - raising water levels upstream reducing effective (intergranular) pressures in the soils
 - scouring of the river bed below the ice which may unload the toe of adjacent slopes
 - raising the level at which ice scours the upstream banks
 - rapid drawdown of upstream water levels when an ice jam breaks up
 - increased scour of river bed and adjacent slopes downstream of an ice jam when it is released
- (ix) South facing slopes tend to thaw and fail at a faster rate than north facing slopes. Material displaced into the river bed may tend to divert the stream channel to the south and cause undercutting of the north facing slope which then becomes unstable because of accelerated oversteepening. This effect is particularly noticeable in the lower reaches of the Mountain River.

INTERPRETATION OF FIELD INVESTIGATIONS

- (i) Multiple retrogressive failures are characteristic of glaciolacustrine sediments and Cretaceous shales in river banks more than 100 feet in height. In some places, at least, the material has failed while in the frozen state. Although the backtilting characteristic gives the appearance of rotational failures to many of these slides, a combination translational-rotational mode of failure is indicated.
- (ii) Slope failures inland from river valleys are not common south of the Fort Good Hope area. Where they are encountered inland, failure has usually been precipitated by forest fire damage to the organic cover. In such instances flows on inclines as low as 4° were observed.

(iii) Active layer failures usually exhibit a high degree of deformation or plasticity and hence fall in the category of "flows".

(iv) Most stream tributaries to the Mackenzie River have the characteristic failure pattern of meandering streams, with unstable slopes on the outside of meander loops. When fire has damaged the organic cover, however, failing banks may be continuous.

(v) The oversteepening of river banks by the abrading action of river ice during spring break-up is considered a major contributing factor in the initiation of landslides.

(vi) Slopes seem to stabilize at about 15° for most soil types although this figure is probably higher for glacial tills. The angle of repose for granular soils is approximately 35° , so that failing soil slopes are usually inclined in the 15 - 35° range. Slopes steeper than 35° generally involve pre-Quaternary materials which have higher cohesion, or Quaternary soils which are temporarily oversteepened and unstable.

ENGINEERING SIGNIFICANCE OF MASS WASTING IN THE MACKENZIE VALLEY

In addition to the usual considerations associated with the siting and design of river crossings, attention should be given to certain factors related to the presence of permafrost and severe climatic conditions in the Mackenzie Valley.

(i) In any location where the organic cover has been damaged or destroyed by forest fires, failures will be initiated on sloping ground and the rate of failure of previously failing slopes will be accelerated. Inland slopes which are otherwise stable will be vulnerable to failure under such conditions.

(ii) Sharp river bends should be avoided as crossing points; meandering streams should be crossed between meander bends.

(iii) As crossing sites are considered in terms of their potential stability, indications of any recent active movement of the slope should not be construed as evidence that a particular slope has been stabilized. Many types of failures are retrogressive and may continue to move inland for thousands of feet.

(iv) Failure planes associated with multiple retrogressive slides are quite deep compared to normal depths used in pipeline trenching operations. Hence movements of blocks of ground in this fashion could displace entire sections of pipe. Since the pipe would be buried below the active layer shallow, active layer slides would be more likely to result in eventual exposure of the pipe and its subsequent settlement.

(v) Support structures for over-water crossings on the tributaries should be designed not only to withstand pressures of ice moving downstream but to withstand pressures of ice backed up from the Mackenzie River when water levels are raised by ice jamming.

(vi) Underwater crossings must be designed for depths of river bed scour which are greater than for normal flood conditions. The effects of ice jamming described earlier, and of ice scour due to high ice loads carried by rivers such as the Great Bear River should be considered.

REQUIREMENTS FOR FURTHER STUDY

The following aspects of slope stability in the Mackenzie Valley require further and more intensive study.

- (i) The mechanics of failure of frozen soils in permafrost terrain.
- (ii) The significance of seismicity in the initiation of slope failures.
- (iii) The relationship between the occurrence of slope failures and annual and seasonal climatic variations.
- (iv) Rates of regression of failing slopes.

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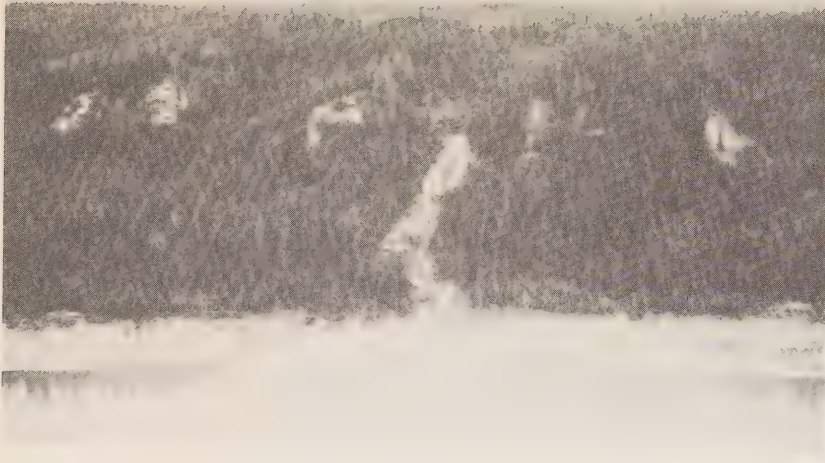


Figure 1 Earthflow, elongate (Qa) Mackenzie River, north of Wrigley.

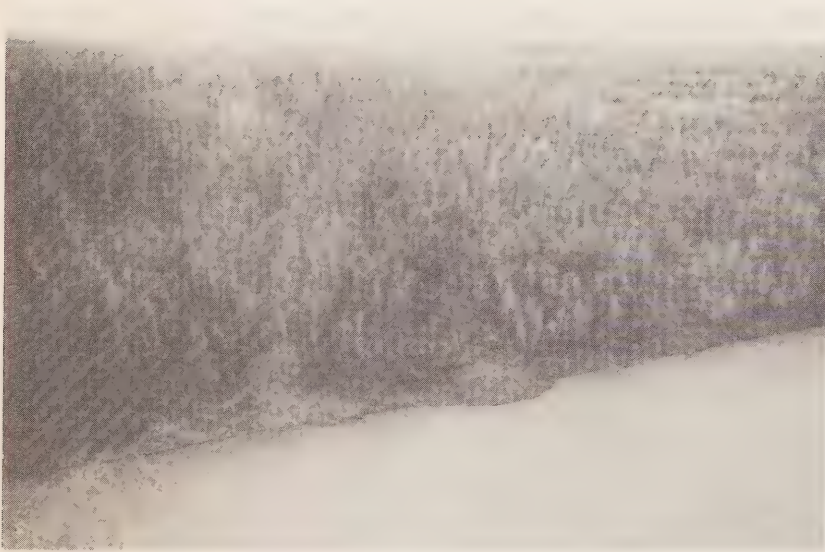


Figure 2 Earthflow, lobate (Qa) Mackenzie River, north of Fort Simpson.



Figure 3 Earthflows in recently burnt area west of Wrigley (Qa).



Figure 4 Closer view of Figure 3.



Figure 5 Detachment slides in granular soil (Qa) Mackenzie River near Blackwater River.

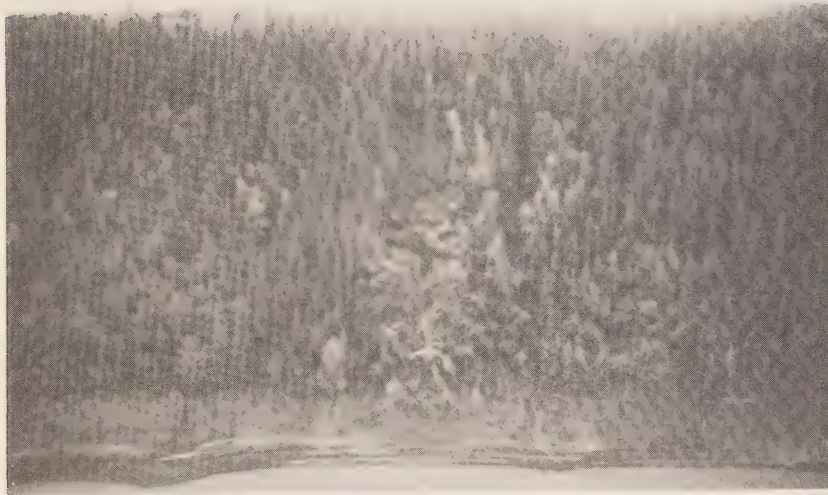


Figure 6 Detachment slide in Cretaceous shale (Ka).



Figure 7 Large scale retrogressive slide; (Q1) Mackenzie River near Root River.



Figure 8 Large scale retrogressive slide; (Q1) Mackenzie River near Root River.

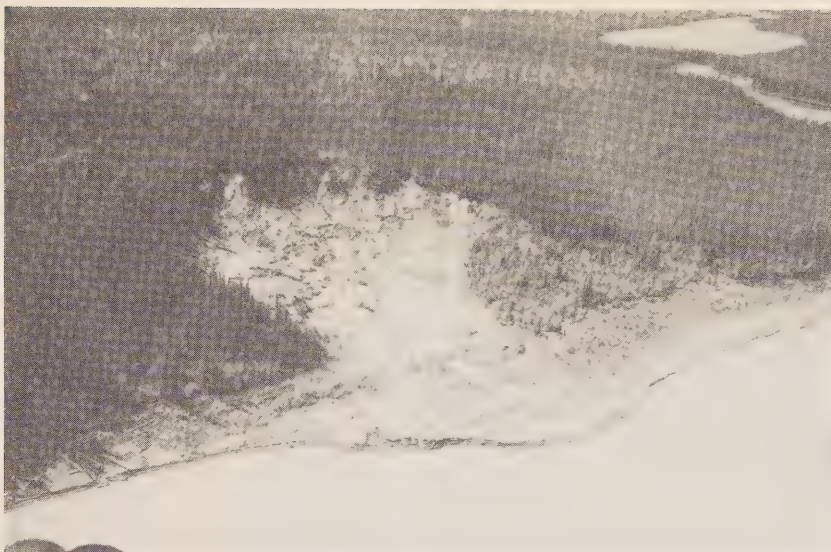


Figure 9 Large scale retrogressive failure; Quaternary sediments (Q1) Mackenzie River opposite Fort Norman.



Figure 10 Large scale retrogressive slides, Quaternary sediments (Q1) Mackenzie River south of Fort Norman.

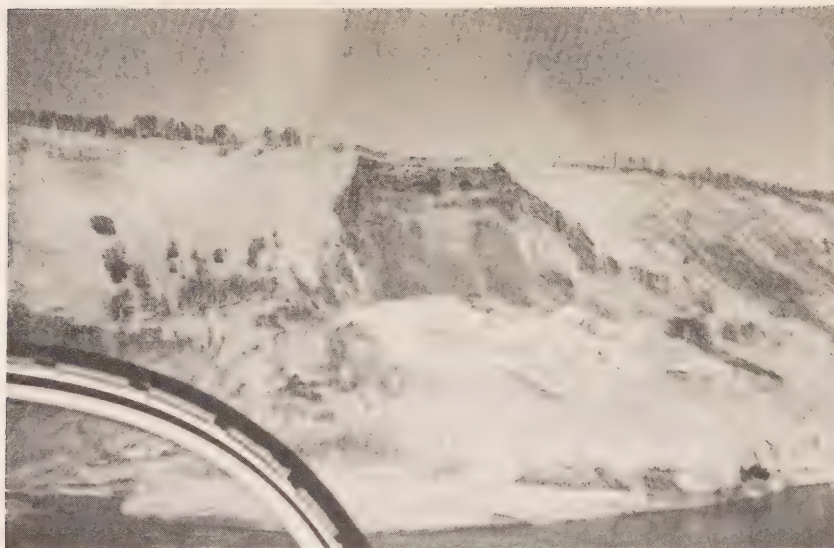


Figure 11 Large scale retrogressive failure in Cretaceous shale (K1) Arctic Red River.

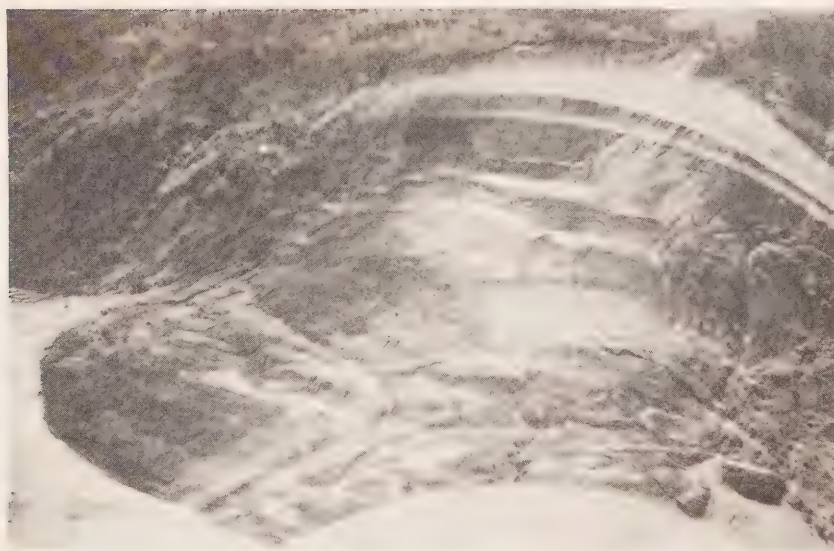


Figure 12 Large scale retrogressive failure in Cretaceous shale (K1) Mountain River.

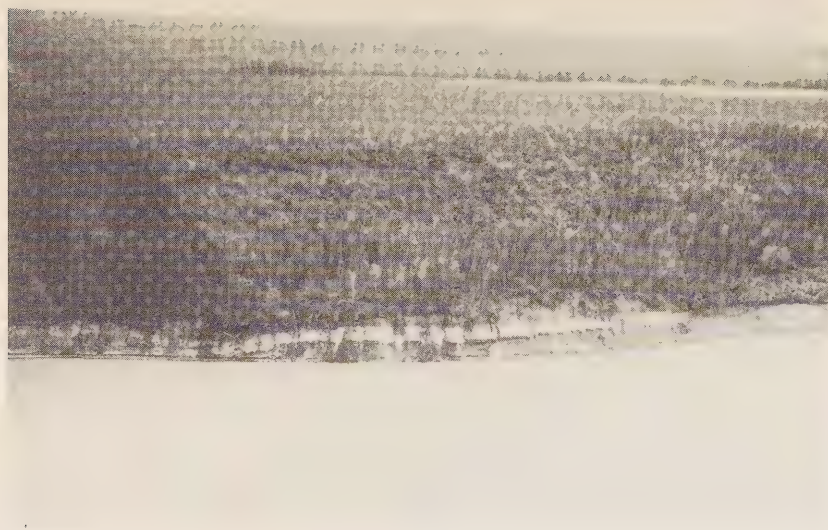


Figure 13 Large scale slumping in Tertiary sediments (Te)
Mackenzie River south of Fort Norman.

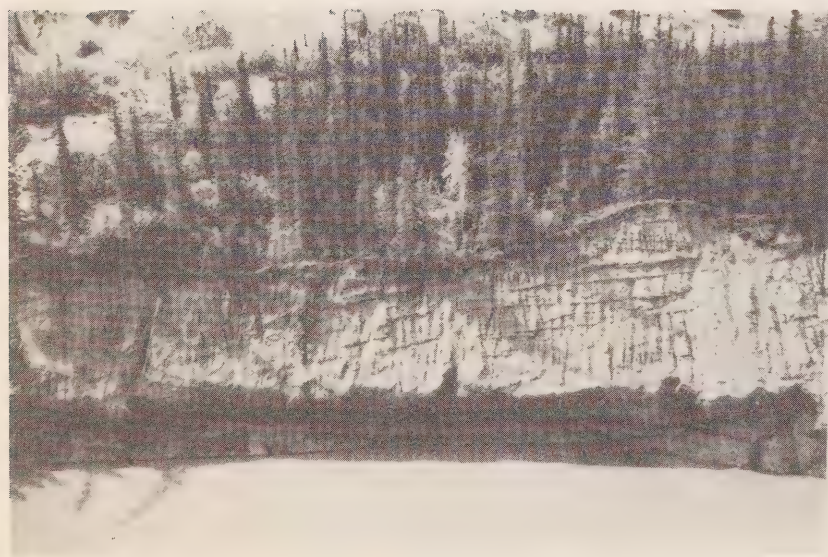


Figure 14 Undercutting of detached frozen slumped block (Q1)
Mackenzie River north of Sans Sault Rapids.



Figure 15 River bank in Devonian sediments (D) Liard River.



Figure 16 Slump in Devonian sediments (D) Liard River.



Figure 17 Large scale multiple retrogressive slide (Q1)
Mackenzie River north of Fort Simpson.



Figure 18 Gullying in glacial till (Qa) Mackenzie River
south of Fort Simpson.

TABLE 1 - RIVER BANK STABILITY MAPS

GEOLOGIC AGE	DESCRIPTION	LEGEND		TYPICAL SLOPE CHARACTERISTICS
		MAP NOTATION	MODE OF EROSION	
Quaternary and Recent	Granular and fine grained (cohesive) uncemented clastic sediments. (Soil cover)	Qs	Negligible, some mass transport of beach and lower slope material by water and river ice.	Stable slopes, vegetated, usually 15° or less. Burnt areas unstable at 5° or less.
		Qa	Mass movement confined to active layer. Failures also shallow in non-permafrost areas. Mainly earthflows, detachment slides, solifluction. Gully erosion and slope wash.	Slope angle 15-35°. Displaced material usually highly deformed due to high moisture contents in active layer. Slopes usually less than 100 feet high.
		Ql	Large scale retrogressive failures (translational slides, slumps, flows); usually accompanied by large scale gullying. Characteristic of glaciolacustrine sediments overlain by glacioluvial sands.	Steep slopes greater than 100 feet in height. Displaced blocks usually relatively undeformed during movement; sometimes consist of frozen soil and often exhibit backward tilt.
Tertiary	Weakly cemented mainly clastic sediments - sandstone, limestone, conglomerate, shale.	Te	Gullying, slope wash, infrequent slumping.	Moderate to steep upper slope; talus accumulation at toe consisting of granular and fragmented rock debris.
		Ka	Gullying, slope wash, shallow active layer slides.	Bank height less than 100 feet. Weathered slopes generally less than 35°.
		Kl	Large scale retrogressive failures of high shale banks.	Steep shale banks unstable at heights of over 100 feet. Low shale content slopes are less susceptible to slumping.
Cretaceous	Weak soft shale; weakly cemented sandstone and siltstone.	K	Undifferentiated	
		D	Rockfalls, infrequent slumping. Some high shale content banks more susceptible to gullying and slumping.	Resistant rocks form steep upper valley walls, flatter talus accumulation at toe. Softer shales erode to low angle valley walls (< 35°).
Devonian	Mainly well cemented, resistant sedimentary rock. Limestone, sandstone, dolomite shale.			



RIVER BANK STABILITY MAPS

CANADA SHEET 95 H



GSC MAP 3-1973

RIVER BANK STABILITY MAP

To accompany
THE STABILITY OF NATURAL SLOPES IN THE MACKENZIE VALLEY
by J. A. Code

Environmental Social Program Report 73-9

Prepared by the Department of Energy, Mines and Resources
for the Environmental-Social Program, Northern Pipelines

LEGEND

GEOLOGIC AGE	DESCRIPTION	MAP NOTATION	MODE OF EROSION	TYPICAL SLOPE CHARACTERISTICS
Quaternary and Recent	Granular and fine grained (cohesives) uncemented clastic sediments (Soil cover)	Qs	Negligible: some mass transport of beach and lower slope material by water and river ice	Stable slopes, vegetated, usually 15° or less. Burnt areas unstable at 9° or less
		Qa	Mass movement confined to active layer. Failures also shallow in non-permafrost areas. Mainly earthflows, disintegrated slides, solifluction. Gully erosion and slope-wash	Slope angle 15-35°. Displaced material usually highly deformed due to high moisture contents in active layer. Slopes usually less than 100 feet high
		Ql	Large scale retrogressive failures (translational slides, slumps, flows), usually accompanied by large-scale gullying. Characteristic of glacio-lacustrine sediments, overlain by glaciofluvial sands	Steep slopes greater than 100 feet in height. Displaced blocks usually relatively undeformed during movement. Sometimes consist of loam soil and often exhibit backward tilt
Tertiary	Weakly cemented, mainly clastic sedimentary: sandstone, limestone, conglomerate, shale	Te	Gullying, slope wash, infrequent slumping	Moderate to steep upper slope, talus accumulation at toe consisting of granular and fragmented rock debris
Cretaceous	Weak soft shale, weakly cemented sandstone and siltstone	Ka	Gullying, slope wash, shallow active layer slides	Bank height less than 100 feet. Shallow active layer slides. Weathered slopes generally less than 35°
		Kl	Large scale retrogressive failures of high shale banks	Steep shale banks unstable at heights of over 100 feet. Low shale content slopes are less susceptible to slumping
Devonian	Mainly well cemented, resistant sedimentary rock: limestone, sandstone, dolomite, shale	D	Undifferentiated	Resistant rocks form steep upper valley walls, flatter talus accumulation at toe. Shale shales erode to low angle valley walls (< 35°)

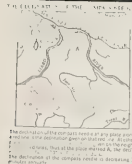
NOTES

- Vertical sequences of the above units observed in the field are shown with components divided by horizontal stroke. For example Qs denotes Quaternary with shallow slides over Cretaceous sandstone. Thicknesses of units not measured.
- Notations showing combinations of above subdivisions such as Qas indicate predominance of Qa with subordinate Qs.
- Transitions between units are often gradual rather than abrupt, in such instances boundaries are shown arbitrarily.
- Where the above notation is applied to meandering rivers, instability if indicated applies only to outside banks of meander loops.

Compiled by J. A. Code from information collected in 1971-1972

Cartography by Geological Survey of Canada

Printed by Surveys and Mapping Branch 1973



REFERENCE

Rock	Symbol	Reference
Bedrock (unconsolidated)	Qs	Quaternary
Bedrock (consolidated)	Qa	Quaternary
Bedrock (consolidated)	Ql	Quaternary
Bedrock (consolidated)	Te	Tertiary
Bedrock (consolidated)	Ka	Cretaceous
Bedrock (consolidated)	Kl	Cretaceous
Bedrock (consolidated)	D	Devonian

FORT SIMPSON
NORTHWEST TERRITORIES
DISTRICT OF MACKENZIE
Scale 1:250,000
1 inch = 4 miles approximately

REFERENCE

Symbol	Reference
Qs	Quaternary
Qa	Quaternary
Ql	Quaternary
Te	Tertiary
Ka	Cretaceous
Kl	Cretaceous
D	Devonian



GSC MAP 5-1973
RIVER BANK STABILITY MAP

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by J.A. Code

Environmental Social Program Report 73-9

Prepared by the Department of Energy, Mines and Resources
for the Environmental-Social Program, Northern Pipelines

LEGEND				
GEOLOGIC AGE	DESCRIPTION	MAP NOTATION	MODE OF EROSION	TYPICAL SLOPE CHARACTERISTICS
Quaternary and Recent	Granular and fine grained (cohesionless) unconsolidated clastic sediments (Silt, sand)	Qs	Negligible, some mass transport of beach and lower slope material by water and river ice	Stable slopes, vegetated, usually 15° or less. Buried areas unstable at 5° or less
		Qa	Mass movement confined to active layer. Failures also shallow in non-permafrost areas. Mostly earthflows, detachment slides, solifluction. Gully erosion and slope-wash	Slope angle 15-35°. Displaced material usually highly deformed due to high moisture contents in active layer. Slopes usually less than 100 feet high
		Qi	Large scale, retrogressive failures (translational slides, slumps, flows) usually accompanied by large scale gullying. Characteristic of glaciolacustrine sediments overlain by glacioluvial sands	Steep slopes greater than 100 feet in height. Displaced blocks usually relatively undeformed during movement. Sometimes consist of frozen sand and often exhibit backward tilt
Tertiary	Weakly cemented mainly clastic sediments: sandstone, limestone, conglomerate, shale	Te	Gullying, slope wash, infrequent slumping	Moderate to steep upper slope, talus accumulation at toe consisting of granular and fragmented rock debris
		Ka	Gullying, slope wash, shallow active layer slides	Bank height less than 100 feet. Weathered slopes generally less than 35°
Cretaceous	Weak soft shale, weakly cemented sandstone and siltstone	Ki	Large scale retrogressive failures of high shale banks	Steep shale banks unstable at heights of over 100 feet. Low shale content slopes are less susceptible to slumping
		K	Undifferentiated	
Devonian	Mainly well cemented, resistant sedimentary rock: Limestone, sandstone, dolomite, shale	D	Rockfalls, infrequent slumping. Some high shale content banks more susceptible to gullying and slumping	Resistant rocks form steep upper valley walls. Flatter talus accumulation at toe. Softer shales erode to low angle valley walls (< 35°)

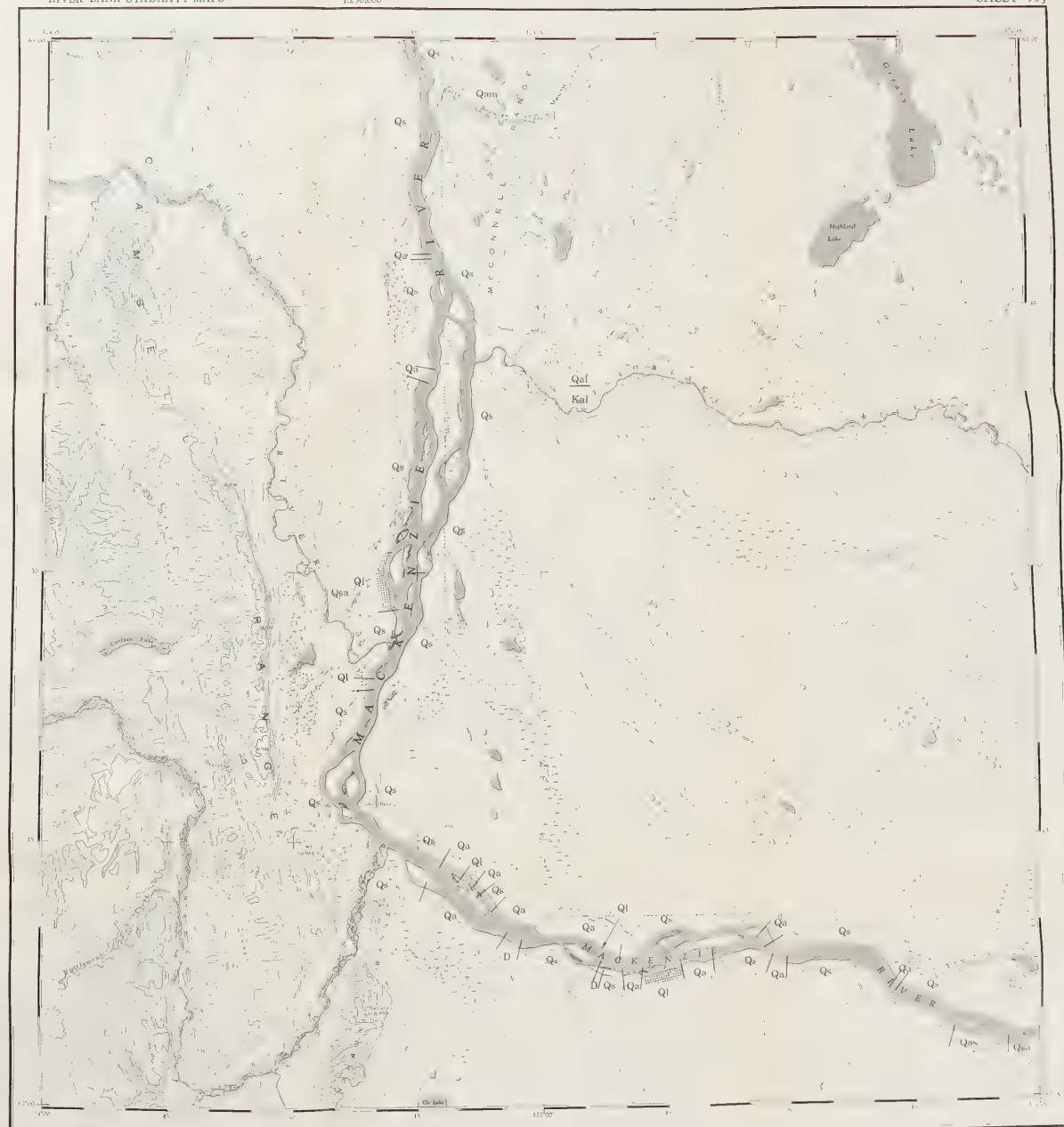
NOTES

- Vertical sequences of the above units observed in the field are shown with components divided by horizontal stroke. For example $\frac{Qs}{Qa}$ denotes Quaternary with shallow slides over Cretaceous sediments. Thicknesses of units not measured.
- Notations showing combinations of above subdivisions such as Qa_i indicate predominance of Qa with subordinate Qi .
- Transitions between units are often gradual rather than abrupt. In such instances boundaries are chosen arbitrarily.
- Where the above notation is applied to meandering rivers, instability if indicated applies only to outside banks of meander loops.

Compiled by J.A. Code from information collected in 1971-1972

Cartography by Geological Survey of Canada

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THE LOCATION OF THIS SHEET IN THE NORTHWEST TERRITORIES

Geological Survey of Canada

CAMSELL BEND
NORTHWEST TERRITORIES
DISTRICT OF MACKENZIE

Scale 1:250,000

1 inch = 4 Miles Approximately



REFERENCE

- Weight in water scale
- Time in water
- W. on or in water
- Bottom in water
- R.C. or P. on
- Recessed or in water
- At bottom of water
- Bottom in water

REFERENCE

see also Oil Development Committee
Committee Publications



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by J.A. Code

Environmental Social Program Report 73-9

Prepared by the Department of Energy, Mines and Resources
for the Environmental-Social Program, Northern Pipelines

LEGEND				
GEOLOGIC AGE	DESCRIPTION	MAP NOTATION	MODE OF EROSION	TYPICAL SLOPE CHARACTERISTICS
Quaternary and Recent	Granular and fine grained (cohesive) unconsolidated clastic sediments (Soil cover)	Qs	Negligible, some mass transport of beach and lower slope material by water and river ice	Stable slopes, vegetated usually 15° or less. Barren areas unstable at 5° or less
		Qa	Mass movement confined to active layer. Failures also shallow in non-permafrost areas. Many earthflows, dissection, slides, solifluction. Gully erosion and slope wash	Slope angle 15-35°. Displaced material usually highly deformed due to high moisture content at active layer. Slopes usually less than 100 feet high
		Qi	Large scale, retrogressive failures (translational slides, slumps, flows), usually accompanied by large scale gullying. Characteristic of glacio-lacustrine sediments overtaken by glaciofluvial sands	Steep slopes greater than 100 feet in height. Displaced blocks usually relatively undeformed during movement. Sometimes consist of frozen soil and often exhibit backward tilt
Tertiary	Weakly cemented, mainly clastic sediments: sandstone, limestone, conglomerate, shale	Te	Gullying, slope wash, infrequent slumping	Moderate to steep upper slope; talus accumulation at toe consisting of granular and fragmented rock debris
Cretaceous	Weak soft shale; weakly cemented sandstone and siltstone	Ka	Gullying, slope wash, shallow active layer slides	Bank height less than 100 feet. Weathered slopes generally less than 35°
		Ki	Large scale retrogressive failures of high shale banks	Steep shale banks unstable at heights of over 100 feet. Low shale content slopes are less susceptible to slumping
Devonian	Mainly well cemented, resistant sedimentary rock: limestone, sandstone, dolomite, shale	K	Undifferentiated	
		D	Rock falls, infrequent slumping. Some high shale content banks more susceptible to gullying and slumping	Resistant rocks form steep upper valley walls, flatter talus accumulation at toe. Softer shales erode to low angle valley walls (< 35°)

- NOTES
- Vertical sequence of the above units observed in the field are shown with components divided by horizontal strokes. For example, Qs denotes Quaternary with shallow slides over Cretaceous sediments. Thickness of units not measured.
 - Notations showing combinations of above subdivisions such as Qas indicate predominance of Qa with subordinate Qs.
 - Transitions between units are often gradual rather than abrupt; in such instances boundaries are chosen arbitrarily.
 - Where the above notation is applied to meandering rivers, instability it indicates applies only to outside banks of meander loops.

Compiled by J.A. Code from information collected in 1971-1972

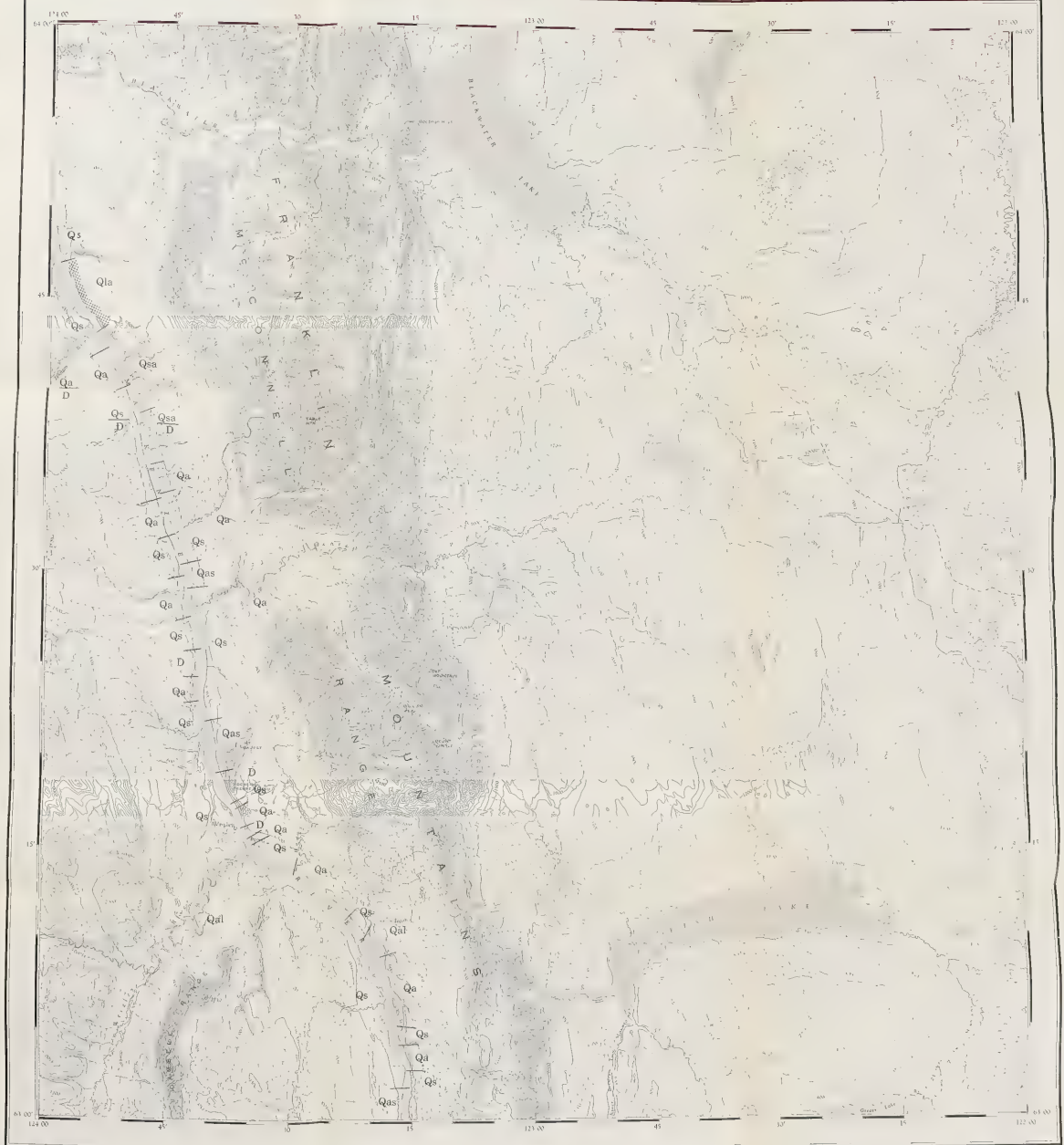
Cartography by Geological Survey of Canada

Printed by Surveys and Mapping Branch 1973

RIVER BANK STABILITY MAPS

1:250,000

SHEET 95-O



THE OUTLINES OF THE MAPS NEEDLE 1-1

Business Monitor Project 55

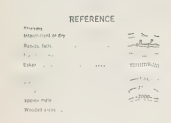
Control Station 55-1-1
Photocopy of the Survey of Canada
Map Series 1:250,000



WRIGLEY
NORTHWEST TERRITORIES
DISTRICT OF MACKENZIE
Scale 1:250,000
1 inch = 4 Miles Approximately



REFERENCE





GSC MAP 8-1973
RIVER BANK STABILITY MAP

To accompany
THE STABILITY OF NATURAL SLOPES IN THE MACKENZIE VALLEY
by J.A. Code

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Prepared by the Department of Energy, Mines and Resources
for the Environmental-Social Program, Northern Pipelines

LEGEND			
GEOLOGIC AGE	DESCRIPTION	MAP NOTATION	TYPICAL SLOPE CHARACTERISTICS
Quaternary and Recent		Qs	Negligible, some mass transport of beach and lower slope material by water and over ice
	Granular and fine grained (cohesive) unconsolidated clastic sediments (Soil cover)	Qa	Mass movement confined to active layer. Failures also shallow in non-permafrost areas. Mainly periglacial detachment, slumps, solifluction. Gully erosion and slope-wash
		Ql	Large scale, retrogressive failures in transitional slides, slumps, flows, usually accompanied by large scale gullying. Characteristic of glaciolacustrine sediments overlain by glaciolacustrine sands
Tertiary	Weakly cemented, mainly clastic sediments: sandstone, limestone, conglomerate, shale	Te	Gullying, slope wash, infrequent slumping. Moderate to steep upper slope, talus accumulation at toe consisting of granular and fragmented rock debris
Cretaceous	Weak, soft shale; weakly cemented sandstone and siltstone	Ka	Gullying, slope wash, shallow active layer slides
		Kl	Large scale retrogressive failures of high shale banks. Low shaly content slopes are less susceptible to slumping
Devonian	Mainly well cemented, resistant sedimentary rock: Limestone, sandstone, dolomite, shale	D	Rockfalls, infrequent slumping. Some high shale content banks more susceptible to gullying and slumping. Resistant rocks form steep upper valley walls, lateral talus accumulation at toe. Softer shales erode to low angle valley walls (< 35°)

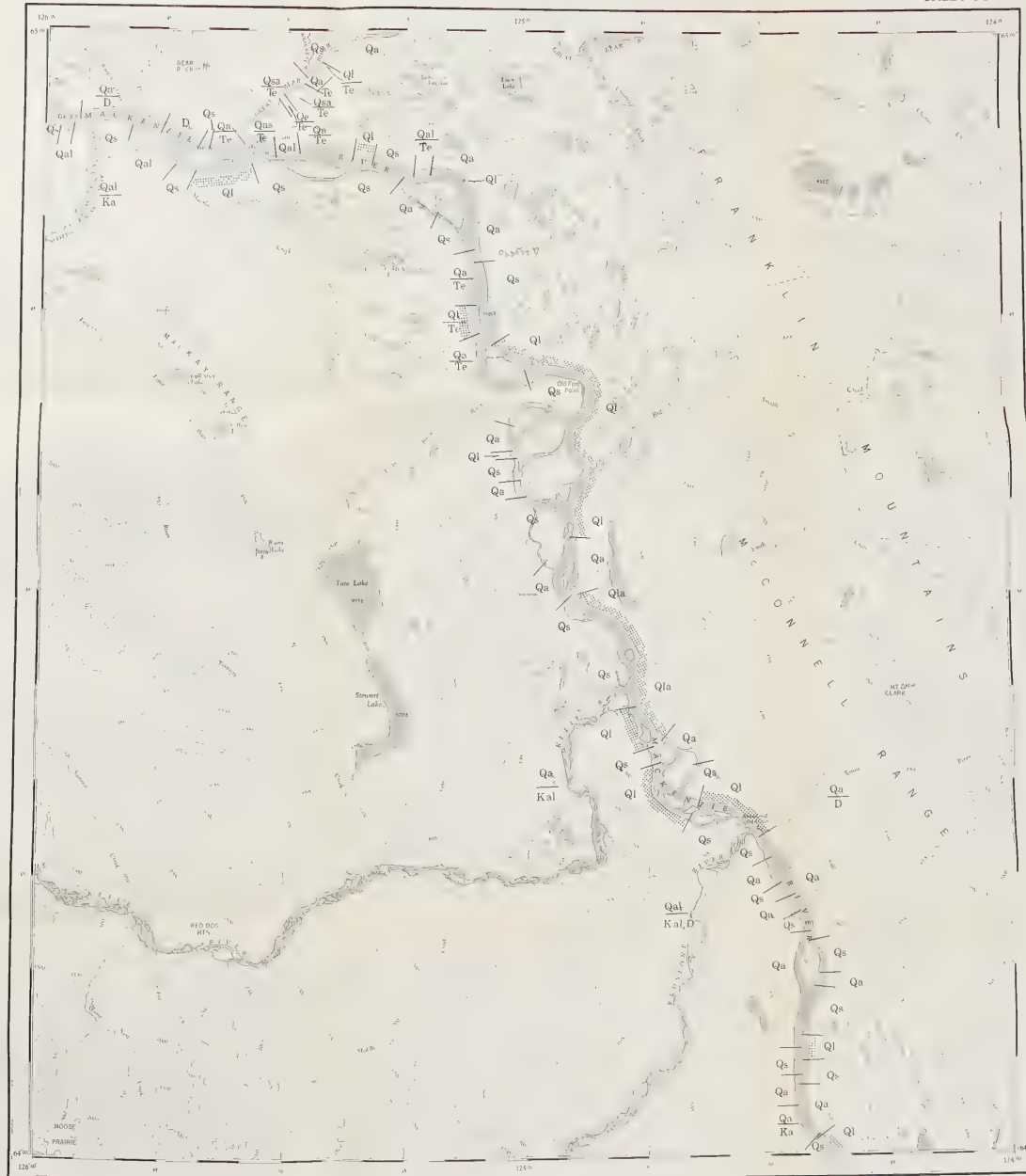
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- Vertical sequences of the above units observed in the field are shown with components divided by horizontal stroke. For example $\frac{Qa}{Te}$ denotes Quaternary with shallow slides over Cretaceous sediments. Thicknesses of units not measured.
- Notations showing combinations of above subdivisions such as Qas indicate predominance of Qa with subordinate Qs .
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- Where the above notation is applied to meandering rivers, instability indicated applies only to outside banks of meander loops.

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Printed by Surveys and Mapping Branch 1973



DECLINATION OF THE COMPASS NEEDLE 1957



The declination of the compass needle is shown along a line with the declination given in degrees and minutes. At each place the declination is given in degrees and minutes. The declination is given in degrees and minutes. The declination is given in degrees and minutes.

Surveyed, 1951. Compiled, drawn and printed by the
ARMY SURVEY ESTABLISHMENT S.C.E., 1958-59
Aerial photography by the R.C.A.F., 1958

REFERENCE

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RIVER BANK STABILITY MAP

To accompany
THE STABILITY OF NATURAL SLOPES IN THE MACKENZIE VALLEY
by J A Code

Environmental Social Program Report 73-9

Prepared by the Department of Energy, Mines and Resources
for the Environmental-Social Program, Northern Pipelines

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		Cs	Negligible some mass transport of beach and lower slope material by water and root ice	Shale slopes vegetated usually 15° or less. Bunt areas unstable at 5° or less.
Quaternary and Recent	Granular and fine grained (colloidal) unconsolidated clastic sediments (Belt cover)	Ga	Mass movement continued to active layer. Failures also show in non-permafrost areas. Mainly earthflows, detachment slides, solifluction. Gully erosion and slope-ward	Slope angle 15-35°. Displaced material usually highly deformed due to high temperatures in active layers. Slopes usually less than 100 feet high.
		Ql	Large scale, retrogressive failures (translational slides, slumps, flows), usually accompanied by large scale gullying. Characteristics of glacio-lacustrine sediments overlain by glaciofluvial sands	Steep slopes greater than 100 feet in height. Displaced blocks usually relatively undeformed. During movement sometimes consist of frozen soil and often extend backward tilt.
Tertiary	Weakly cemented mainly clastic sediments-sandstone, limonstone, conglomerate shale	Te	Gullying, slope slumping, erosion, wash	Moderate to steep upper slope with accumulation at toe consisting of granular and fragmented rock debris.
Crataceous	Weak soft shale, weakly cemented sandstone and siltstone	Ka	Gullying, slope wash	Bank height less than 100 feet. Weighted slopes generally less than 35°.
		Kl	Large scale retrogressive failures of high shale banks	Steep slope banks unstable at heights of over 100 feet. Low shale convex slopes are less susceptible to slumping.
		K	Undifferentiated	
Devonian	Mainly well cemented, resistant sedimentary rock. Limonstone, sandstone, dolomite shale	D	Rockfalls, retrogressive slumps. Some high shale convex banks more susceptible to gullying and slumping	Resistant rock forms steep upper valley walls, rather stable accumulation at toe. Softer shales erode to low angle valley floor (< 35°).

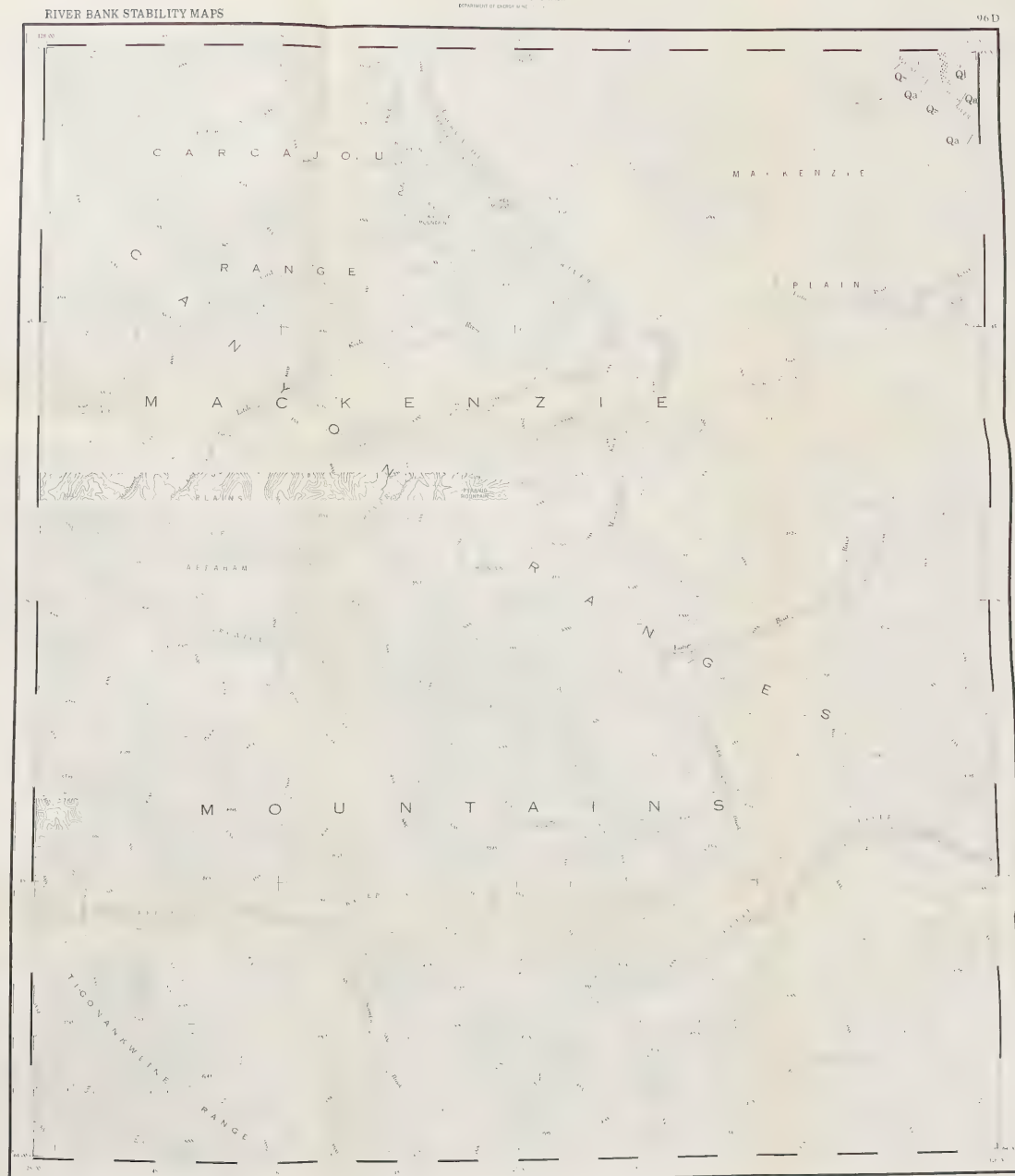
NOTES

1. Vertical sequences of the above units observed in the field are shown with components divided by horizontal stroke. For example, $\frac{Q_2}{K}$ denotes Quaternary with shallow slides over Cretaceous sediments. Thicknesses of units not measured.
2. Notations showing combinations of above subdivisions such as Q_2s indicate predominance of Q_2 with subordinate Q_s .
3. Transitions between units are often gradual rather than abrupt, in such instances boundaries are chosen arbitrarily.
4. Where the above notation is applied to meandering rivers, instability if indicated applies only to outside banks of meander loops.

Compiled by J. A. Code from information collected in 1971-1972

Cartography by Geological Survey of Canada

Printed by Surveys and Mapping Branch 1973



CARCAJOU CANYON

DISTRICT OF MACKENZIE
NORTHWEST TERRITORIES

Scale 1:250,000 Échelle





RIVER BANK STABILITY MAP

To accompany
THE STABILITY OF NATURAL SLOPES IN THE MACKENZIE VALLEY
by J.A. Code

Environmental Social Program Report 73-9

Prepared by the Department of Energy, Mines and Resources
for the Environmental-Social Program, Northern Pipelines

LEGEND

GEOLOGIC AGE	DESCRIPTION	MAP NOTATION	MODE OF EROSION	TYPICAL SLOPE CHARACTERISTICS
Quaternary and Recent		Qs	Negligible, some mass transport of beach and lower slope material by water and river ice	Stable slopes, vegetated, usually 15° or less. Burnt areas unstable at 5° or less
	Granular and fine grained (clayey) unconsolidated clastic sediments (Silt cover)	Qa	Mass movement confined to active layer. Failures also shallow in non-permafrost areas. Mostly earthflows, detachment slides, soilflowing. Gully erosion and slope-wash	Slope angle 15-35°. Displaced material usually highly deformed due to high moisture contents in active layer. Slopes usually less than 100 feet high
		Ql	Large scale, retrogressive failures (translational slides, slump, flows) usually accompanied by large scale gullying. Characteristic of glacio-lacustrine sediments overlain by glaciofluvial sands	Steep slopes greater than 100 feet in height. Displaced blocks usually relatively undeformed during movement, sometimes consist of frozen soil and often exhibit backward tilt
Tertiary	Weakly cemented, mainly clastic sediments: sandstone, limestone, conglomerate and shale	Te	Gullying, slope wash, infrequent slumping	Moderate to steep upper slope talus accumulation at toe consisting of granular and fragmented rock debris
Cretaceous		Ka	Gullying, slope wash, shallow active layer slides	Bank height less than 100 feet. Weathered slopes generally less than 35°
	Weak soft shale, weakly cemented sandstone and siltstone	Kl	Large scale retrogressive failures of high shale banks	Steep shale banks unstable at heights of over 100 feet. Low shale content slopes are less susceptible to slumping
		K	Undifferentiated	
Devonian	Mainly well cemented resistant sedimentary rock: limestone, sandstone, dolomite shale	D	Rockfalls, infrequent slumping. Some high shale content banks more susceptible to gullying and slumping	Resistant rocks form steep upper valley walls. Talus talus accumulation at toe. Softer shales erode to low angle valley walls (< 35°)

NOTES

- Vertical sequences of the above units observed in the field are shown with components divided by horizontal stroke. For example, $\frac{Qa}{Te}$ denotes Quaternary with shallow slides over Cretaceous sediments. Thickness of units not measured.
- Notations showing combinations of above subdivisions such as Qa indicate predominance of Qa with subordinate Qs .
- Transitions between units are often gradual rather than abrupt. In such instances boundaries are chosen arbitrarily.
- Where the above notation is applied to meandering rivers, instability indicated applies only to outside banks of meander loops.

Compiled by J.A. Code from information collected in 1971-1972

Cartography by Geological Survey of Canada

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RIVER BANK STABILITY MAPS



106 H



SANS SAULT RAPIDS NORTHWEST TERRITORIES DISTRICT OF MACKENZIE

Scale 1:250,000 Etchelle

Root of weather: Cône de vent
 River in channel: Rivière en chenal
 Road or bridge: Route ou pont
 River or stream: Rivière ou ruisseau
 Water body: Corps d'eau
 Boundary marked: Frontière marquée

1:100,000

1:50,000

1:25,000

1:12,500

1:6,250

1:3,125

1:1,562

1:781

1:390

1:195

1:97

1:48

1:24

1:12

1:6

1:3

1:1.5

1:0.75

1:0.375

1:0.1875

1:0.09375

1:0.046875

1:0.0234375

1:0.01171875

1:0.005859375

1:0.0029296875

Transverse Meridian Projection
 North American Datum 1983
 Contour Interval 500 feet
 Elevation in feet above Mean Sea Level

Projection Transverse de Mercator
 Méridien géographique central Amérique du Nord, 1983
 Épaisseur des courbes: 500 pieds
 Élévation en pieds au-dessus du niveau moyen de la mer





GSC MAP 12-1973

RIVER BANK STABILITY MAP

To accompany
THE STABILITY OF NATURAL SLOPES IN THE MACKENZIE VALLEY
by J.A. Code

Environmental Social Program Report 73-9

Prepared by the Department of Energy, Mines and Resources
for the Environmental-Social Program, Northern Pipelines

LEGEND				
GEOLOGIC AGE	DESCRIPTION	MAP NOTATION MODE OF EROSION	TYPICAL SLOPE CHARACTERISTICS	
		Qs	Negligible, some mass transport of beach and lower slope material by water and river ice	Stable slopes, vegetated, usually 15° or less. Burnt areas unstable at 5° or less
Quaternary and Recent	Granular and fine grained (cohesive) unconsolidated clastic sediments. (Soil cover)	Ga	Mass movement confined to active layer. Failures also shallow in non-permafrost areas. Many talus piles, detachment ridges, scollification, gully erosion and slope wash	Slope angle 15-35° Displaced material usually highly deformed due to high moisture contents in active layer Slopes usually less than 100 feet high
		Ql	Large scale, retrogressive failures transitional slides, slump, flows, usually accompanied by large scale gullying. Characteristic of glacio-luvial sediments overlain by glaciofluvial sands.	Step slopes greater than 100 feet in height Displaced blocks usually relatively undeformed during movement sometimes consist of frozen soil and often exhibit backward tilt
Tertiary	Weakly cemented mainly clastic sediments-sandstone, limestone, conglomerate, shale	Te	Gullying, slope wash, infrequent slumping.	Moderate to steep upper slopes. Talus accumulation at toe consisting of granular and fragmented rock debris.
Cretaceous	Weak soil shale, weakly cemented sandstone and siltstone	Ka	Gullying, slope wash, shallow active layer slides	Bank heights less than 100 feet. Weard slopes generally less than 35°
		Kl	Large scale retrogressive failures of high shale banks	Steep shale banks unstable at heights of over 100 feet Low shales content slopes less susceptible to slumping
		K	Undifferentiated	
Devonian	Mainly well cemented, resistant sedimentary rock-limestone, sandstone, dolomite shale.	D	Rockfalls, infrequent slumping. Some high talus accumulation. Shales more susceptible to gullying and slumping.	Resistant rocks form steep upper valley walls, flatter talus accumulation at toe. Steep shales erode to low angle valley walls (< 35°)

NOTES

1. Vertical sequences of the above units observed in the field are shown with components divided by horizontal strokes. For example, $\frac{Q_2}{K}$ denotes Quaternary with shallow slides over Cretaceous sediments. Thicknesses of units not measured.
2. Notations showing combinations of above subdivisions such as Q_{21} indicate predominance of Q_2 with subordinate Q_1 .
3. Transitions between units are often gradual rather than abrupt; in such instances boundaries are chosen arbitrarily.
4. Where the above notation is applied to meandering rivers, instability if indicated applies only to outside banks of meander loops.

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Cartography by Geological Survey of Canada

Printed by Surveys and Mapping Branch 1973



FORT GOOD HOPE
NORTHWEST TERRITORIES

Scale 1:250,000
1 Inch to 4 Miles Approximately

10 13

Contour Interval 100 Feet
Elevations in Feet above Mean Sea Level
North American Datum 1927.

32 *Notes*

REFERENCE

Synonym:
clamorous or *dey*
clamorous *leka*
Rapids, falls
March or tramp
Curious:
elevation
depression
approach
Exalt



